Remote Sensing & Photogrammetry

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www.fotogrametria.agh.edu.pl



AKADEMIA GÓRNICZO-HUTNICZA WYDZIAŁ GEODEZJI GÓRNICZEJ I INŻYNIERII ŚRODOWISKA KATEDRA GEOINFORMACJI, FOTOGRAMETRII I TELEDETEKCJI ŚRODOWISKA

Strona Główna <u>Ogłoszenia</u> <u>Zespół</u> <u>Dydaktyka</u> <u>Prace magisterskie</u> Publikacje po 2006 <u>Publikacje do 2006</u> <u>Badania po 2006</u> Badania do 2006



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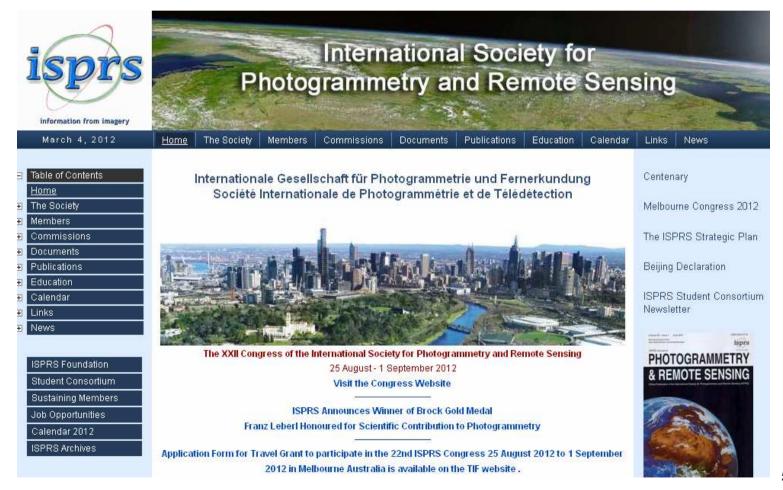
STUDY IN ENGLISH: Photogrammetry and Remote Sensing



Background of remote sensing

- Electromagnetic radiation. Interactions with the atmosphere. Radiation – target interactions.
 Passive and active sensing. Image characteristic (DN, spectral curves).
- Satellite and sensors. Spatial, spectral, radiometric and temporal resolutions. Thermal and microwave imaging. Backgrounds of thermovisin.
- Image interpretation and analysis. Preprocessing. Image enhancement. Image Transformation. Image Classification.

Literature http://www.isprs.org/ Tutorials



http://www.isprs.org/ EducationTutorials

South Africa), by Manos Baltsavias, ETH Zurich: part 1 (1.7 MB 🗟), part 2 (2.1 MB 🗟), part 3 (3.2 MB 🗟)

- Workshop on 'Remote Sensing and GIS for Watershed Managment', by ISPRS Special Interest Group 'Technology Transfer Caravan'
- Extraction of Geospatial Information from High Resolution Optical Satellite Sensors, ISPRS Technical Commission IV Symposium 'Geospatial Databases for Sustainable Development', Goa, India, 27-30 September 2006
- "Space technologies to support the conservation of natural and cultural sites", 25-27 November 2005, Campeche, Mexico

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Tutorials in REMOTE SENSING

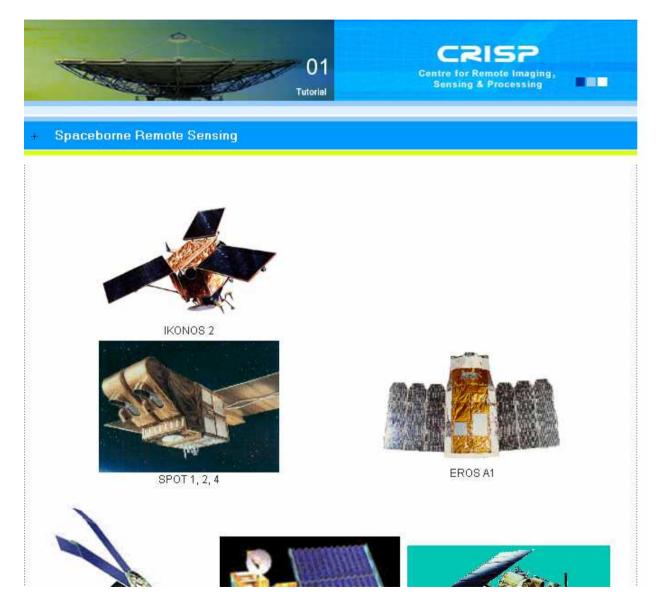
- If The Remote Sensing Core Curriculum, an ASPRS Initiative for Space Age International Education
- Introduction to Remote Sensing, by CSIRO projects
- Canada Centre for Remote Sensing, Education Section
- II The use of Satellite Remote Sensing, CIESIN Thematic Guides
- 🗉 🗳 NASA's Remote sensing Tutorial 🗰
- Remote Sensing and Image Analysis lecture, Berkeley lecture
- Principles of Remote Sensing, Singapore Science Center
- Workshop "Satellite-based Photogrammetry" La Habana, Cuba, 9-13 February 2009

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Tutorials in PHOTOGRAMMETRY

- If Theory of Close Range Photogrammetry, Ch.2 of 'Close Range Photogrammetry and Machine Vision', K.B.Atkinson ed., 1996
- L2 Air Photo Interpretation and Photogrammetry, by International Center for Remote Sensing of Environment
- II Tutorials on optical measurements, by Gigahertz Optik

CRISP



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STATUTES

INTERNATIONAL SOCIETY FOR PHOTOGRAMMETRY AND REMOTE SENSING July, 2008

• STATUTE II - Definitions

 Photogrammetry and Remote Sensing is the art, science, and technology of obtaining reliable information from non-contact imaging and other sensor systems about the Earth and its environment, and other physical objects and processes through recording, measuring, analyzing and representation.

Spatial Information Science is the art, science, and technology of obtaining reliable spatial, spectral and temporal relationships between physical objects, and of processes for integration with other data for analysis, portrayal and representation, independently of scale.

Milestones in the History of Remote Sensing

- 1839 Photography was invented
- 1858 Parisian Photographer, Gaspard Felix Tournachon used a balloon to ascend to a height of 80m to obtain the photograph over Bievre, France
- 1962 The term "Remote Sensing" first appeared
- 1972 The launch of Landsat-1, originally ERTS-1, Remote sensing has been extensively investigated and applied since then
- 1990 Proposed EOS aiming at providing data for global change monitoring. Various sensors have been proposed.
 - Japan's JERS-1 SAR,
 - European ERS Remote Sensing Satellite SAR,
 - Canada's Radarsat
 - Radar and imaging spectrometer data will be the major theme of this decade and probably next decade as well

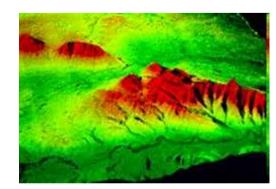
Applications

damage assessment) Forestry (forest cover type discrimination agroforestry mapping, clear cut mapping / regeneration assessment burn delineation , infrastructure mapping / operations support forest inventory, biomass estimation, species inventory) Geology (lithological, structural mapping, sedimentation mapping and monitoringgeo-hazard mapping) Hydrology (flood delineation&mapping, soil mapping) Sea Ice (ice type&concentration, ice motion) Land Cover (Land Cover & Land Use, Land Use Change) Mapping (Planimetry, Digital Elevatin Model, Thematic Mapper) Oceans&Coastal Monitoring

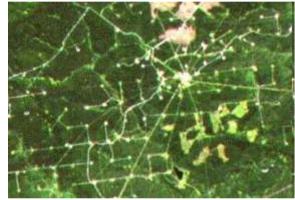
(Ocean Features, Ocean Colour, Oil Spill Detecion)



Agriculture (crop type monitoring,

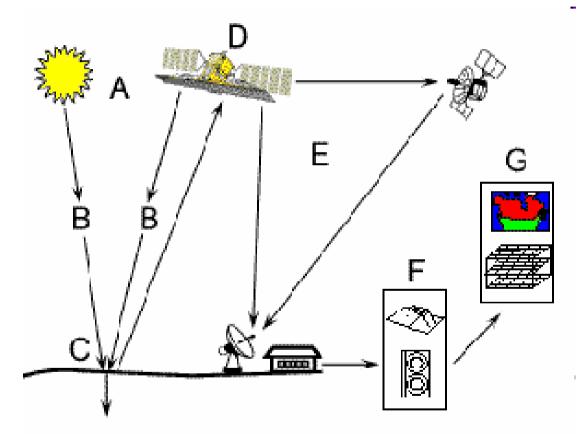








Introduction



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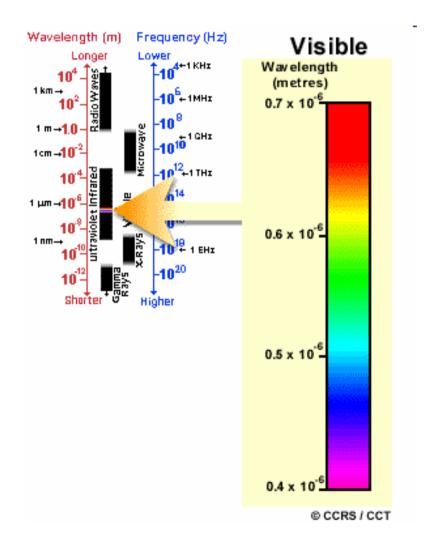
Electromagnetic Energy

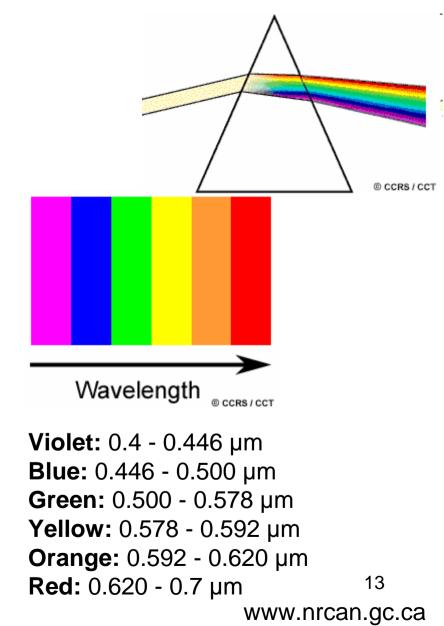
- "Energy is a group of particles travelling through a certain media. Electromagnetic energy is a group of particles with different frequencies travelling at the same velocity. These particles have a dual-mode nature. They are particles but they travel in a wave form.
- Electromagnetic waves obey the following rule:"

$$V = c/\lambda$$

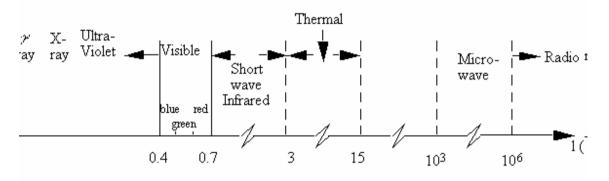
- c : the speed of electromagnetic wave
- λ : wave length
- \mathbf{v} : frequency.

The Electromagnetic Spectrum





Major uses of some spectral wavelength regions



http://www.cnr.berkeley.edu/~gong/textbook/

Wavelength	Use	Wavelength	Use
g ray	Mineral	1.55-1.75 µm	Water content in plant or soil
X ray	Medical	2.04-2.34 µm	Mineral, rock types
Ultraviolet(UV)	Detecting oil spill	10.5-12.5 µm	Surface temperature
0.4-0.45 µm	Water depth, turbidity	3 cm - 15 cm	Surface relief, soil moisture
0.7-1.1 μm	Vegetation vigor	20 cm - 1 m	Canopy penetrat i on, woody biomass

Radiation Laws

The <u>first</u> theory treats electromagnetic radiation as many discrete particles called photons or quanta (terms in Physics).

The energy of a quantum is given by

E=hv

where

- E energy of a quantum (Joules)
- h = 6.626 x 10-34 (Planck's constant)
- v frequency

Energy and wavelegth

Radiation Laws

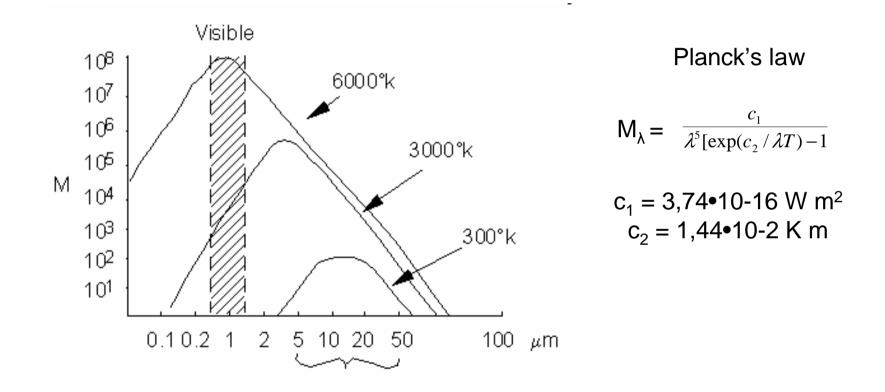
This has implications to remote sensing sensor design. To use the available sensing technology at hand, we will have to balance between wavelength and spatial resolution. If we wish to make our sensor to have higher spatial resolution, we may have to use short wavelength regions.

The **second** radiation theory is Stefan-Boltzmann Law:

$$M = \delta T^4$$

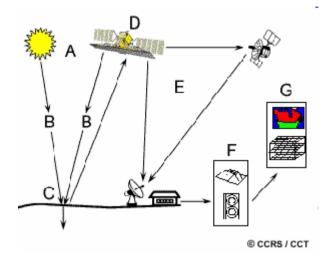
- M: total radiant existence for a surface of a material watts/m²
- δ Stefan-Boltzmann constant, 5.6697 x 10-8 Wm^-2 K^-4
- T: absolute temperature, K

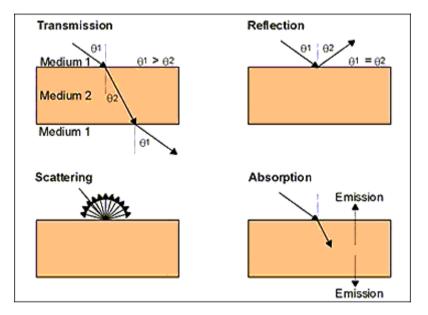
Blackbody radiation



Wien's displacement law

Radiation - Target Interactions



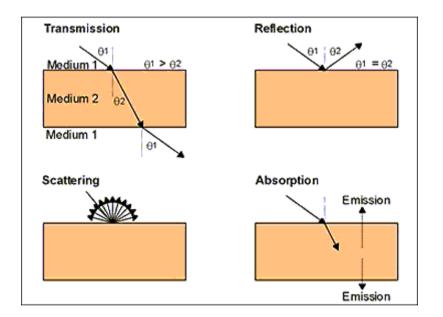


 $1 = \rho + \alpha + \tau$

Real objects

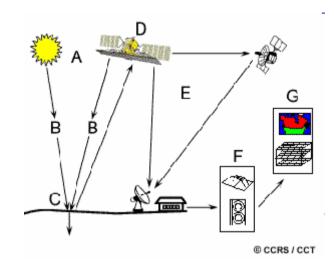
 $M = \sigma \delta T^4$

- M: total radiant existence for a surface of a material watts/m²
- δ Stefan-Boltzmann constant, 5.6697 x 10^-8 Wm^-2 %^-4
- T: absolute temperature,
- σ emmisivity coefficient

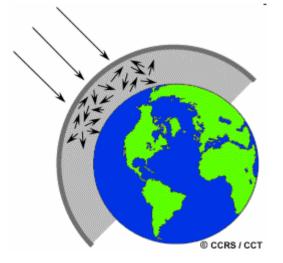


Interactions with the Atmosphere

Scattering occurs when particles or large gas molecules present in the atmosphere interact with and cause the electromagnetic radiation to be redirected from its original path. How much scattering takes place depends on several factors including the wavelength of the radiation, the abundance of particles or gases, and the distance the radiation travels through the atmosphere. There are three (3) types of scattering which take place.



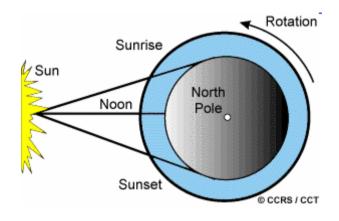




Rayleigh scattering

Occurs when particles are very small compared to the wavelength of the radiation.

These could be particles such as small specks of dust or nitrogen and oxygen molecules. Rayleigh scattering causes shorter wavelengths of energy to be scattered much more than longer wavelengths. Rayleigh scattering is the dominant scattering mechanism in the upper atmosphere. The fact that the sky appears "blue" during the day is because of this phenomenon. As sunlight passes through the atmosphere, the shorter wavelengths (i.e. blue) of the visible spectrum are scattered more than the other (longer) visible wavelengths. At sunrise and sunset the light has to travel farther through the atmosphere than at midday and the scattering of the shorter wavelengths is more complete; this leaves a greater proportion of the longer wavelengths to penetrate the atmosphere.



Mie scattering

Occurs when the particles are just about the same size as the wavelength of the radiation. Dust, pollen, smoke and water vapour are common causes of Mie scattering which tends to affect longer wavelengths than those affected by Rayleigh scattering. Mie scattering occurs mostly in the lower portions of the atmosphere where larger particles are more abundant, and dominates when cloud conditions are overcast.



The final scattering mechanism of importance is called **nonselective scattering**. This occurs when the particles are much larger than the wavelength of the radiation. Water droplets and large dust particles can cause this type of scattering. Nonselective scattering gets its name from the fact that all wavelengths are scattered about equally. This type of scattering causes fog and clouds to appear white to our eyes because blue, green, and red light are all scattered in approximately equal quantities (blue+green+red light = white light).

Absorption

- Is the **other main** mechanism at work when electromagnetic radiation interacts with the atmosphere. In contrast to scattering, this phenomenon causes molecules in the atmosphere to absorb energy at various wavelengths. Ozone, carbon dioxide, and water vapour are the three main atmospheric constituents which absorb radiation.
- **Ozone** serves to absorb the harmful (to most living things) ultraviolet radiation from the sun. Without this protective layer in the atmosphere our skin would burn when exposed to sunlight.
- You may have heard **carbon dioxide** referred to as a greenhouse gas. This is because it tends to absorb radiation strongly in the far infrared portion of the spectrum - that area associated with thermal heating - which serves to trap this heat inside the atmosphere. Water vapour in the atmosphere absorbs much of the incoming longwave infrared and shortwave microwave radiation (between 22µm and 1m). The presence of water vapour in the lower atmosphere varies greatly from location to location and at different times of the year. For example, the air mass above a desert would have very little water vapour to absorb energy, while the tropics would have high concentrations of water vapour (i.e. high humidity).

EM radiation

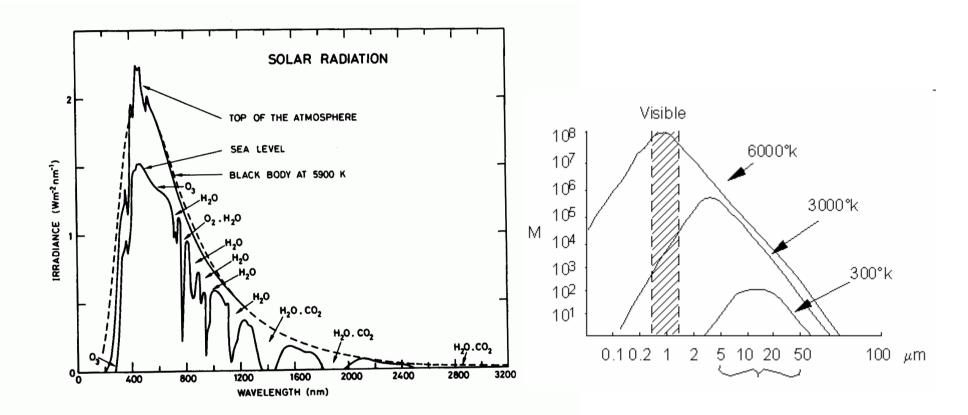
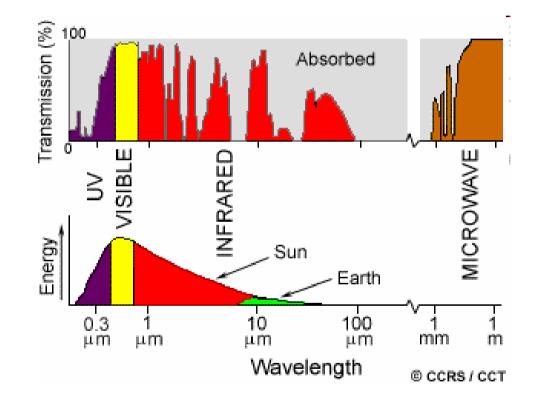
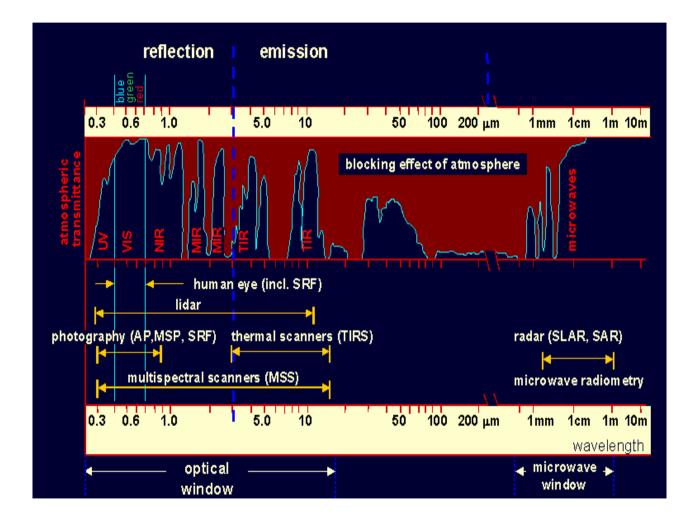


Fig. 4.6. Spectrum of solar radiation (UV, visible, IR) outside the earth's atmosphere and at sea level. (Adapted from Coulson, 1975).

Atmospheric windows

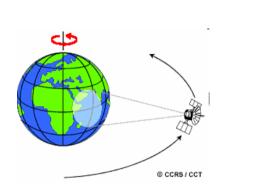


EM radiation

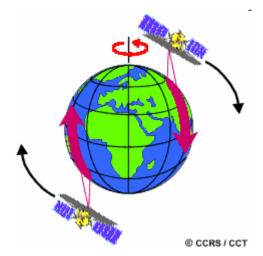


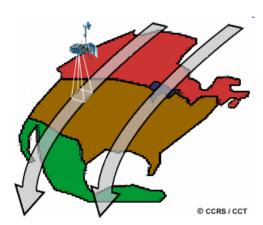
Satellite characteristic

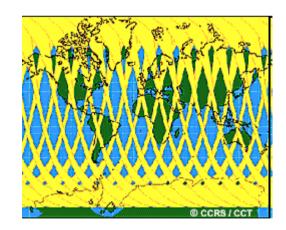
- Geostationary
- Sun-synchronous, local sun time







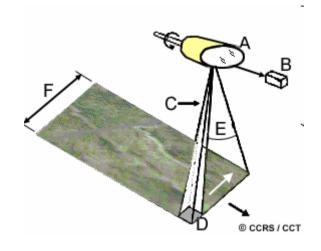


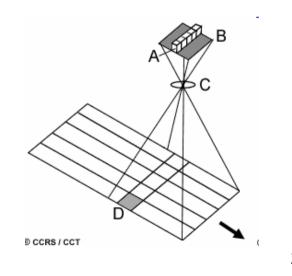


Multispectral Scanning

- multispectral scanner (MSS)
- Across-track scanners
- Along-track scanners, pushbroom scanners

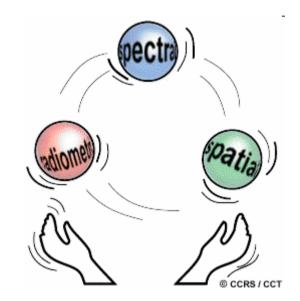
The IFOV (C) of the sensor and the altitude of the platform determine the ground resolution cell viewed (D), and thus the spatial resolution. The angular field of view (E) is the sweep of the mirror, measured in degrees, used to record a scan line, and determines the width of the imaged swath (F).



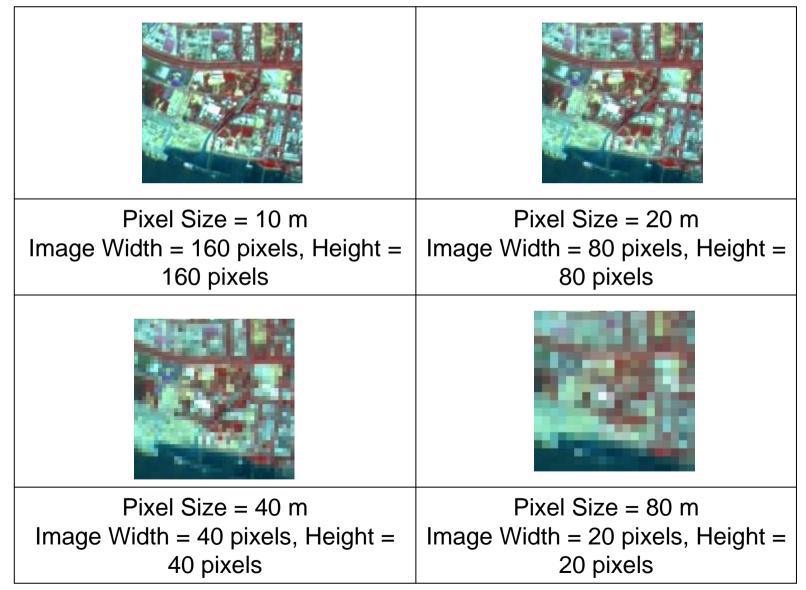


Resolution

- Spatial
- Spectral
- Radiometric
- Temporal



Spatial



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Spatial



Spatial resolution: 0,5; 1,0; 5,0; 30







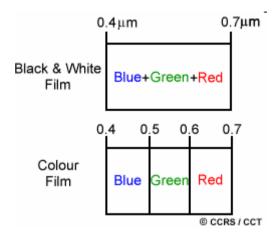
Spatial



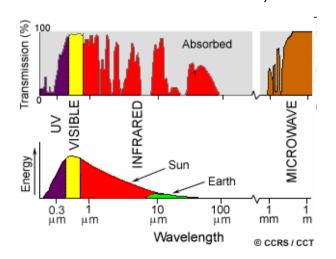


Spectral

- Spectral range (μm):
- UV (0,3-0,4)
- VIS (0,4-0,7)
- VNIR (0,4-1,4)
- IR
- SWIR (1,4-2,5)
- TIR (2,5-14)
- MV (0,75-100cm)



Spectral resolution -PAN 0,3 -MSS 0,1

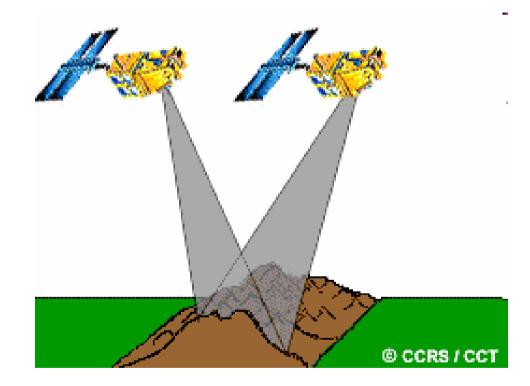


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Radiometric

8-bit quantization (256 levels)	6-bit quantization (64 levels)	
4-bit quantization (16 levels)	3-bit quantization (8 levels)	
2-bit quantization (4 levels)	1-bit quantization (2 levels)	

Temporal



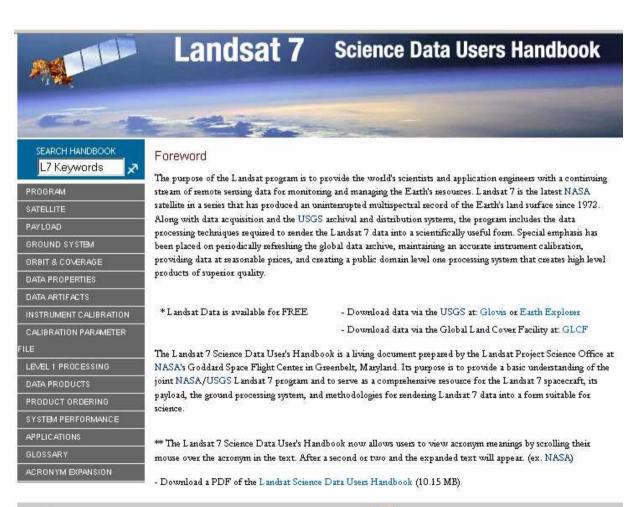
LANDSAT, USA

- The LANDSAT program consists of a series of <u>optical/infrared remote</u> <u>sensing</u> satellites for land observation. The program was first started by The National Aeronautics and Space Administration (NASA) in 1972, then turned over to the National Oceanic and Atmospheric Administration (NOAA) after it became operational. Since 1984, satellite operation and data handling were managed by a commercial company EOSAT. However, all data older than 2 years return to "public domain" and are distributed by the Earth Resource Observation System (EROS) Data Center of the US Geological Servey (USGS).
- The first satellite in the series, LANDSAT-1 (initially named as the Earth Resource Technology Satellite ERTS-1) was launched on 23 July 1972. The satellite had a designed life expectancy of 1 year but it ceased operation only on January 1978. LANDSAT-2 was launched on 22 January 1975 and three additional LANDSAT satellites were launched in 1978, 1982, and 1984 (LANDSAT-3, 4, and 5 respectively). LANDSAT-6 was launched on October 1993 but the satellite failed to obtain orbit. A new satellite LANDSAT-7 was launched in 15 April 1999. Currently, only LANDSAT-5 and 7 are operational.

LANDSAT

					Туре			Sun-Synchronous		
		T 4,5 MSS Sens eristics	or		Altitude			705 km		
		T TM, ETM+ Sei eristics	nsor		Inclinat	ion	98.2 deg			
					Period	Period		99 min		
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Green	2	0.52 - 0.60	30		Band	Wavelength (µ	ım) Resolution (m			
Red	3	0.63 - 0.69	30		Dana	fratololigii (p	,			
Near IR	4	0.76 - 0.90	30	Green	1	0.5 - 0.6		82		
SWIR	5	1.55 - 1.75	30							
Thermal IR	6	10.40 - 12.50	120 (TM) 60 (ETM+)	Red	2	0.6 - 0.7		82		
SWIR	7	2.08 - 2.35	30	Near IR	3	0.7 - 0.8		82		
Panchromatic	;	0.5 - 0.9	15	Near IR	4	0.8 - 1.1		82		

Landsat handbook Landsat Data is available for FREE





+ Privacy Policy and Important Notices



NASA Official: Jim Irons Website Curator: Mike Taylor This page last updated: March 11, 2011

Landsat Data is available for FREE

- Path: 188, Row: 25
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- GEOTIF

	Band	Wavelength (µm)	Resolution (m)
Blue	1	0.45 - 0.52	30
Green	2	0.52 - 0.60	30
Red	3	0.63 - 0.69	30
Near IR	4	0.76 - 0.90	30
SWIR	5	1.55 - 1.75	30
Thermal IR	6	10.40 - 12.50	120 (TM) 60 (ETM+)
SWIR	7	2.08 - 2.35	30
Panchromatic		0.5 - 0.9	15

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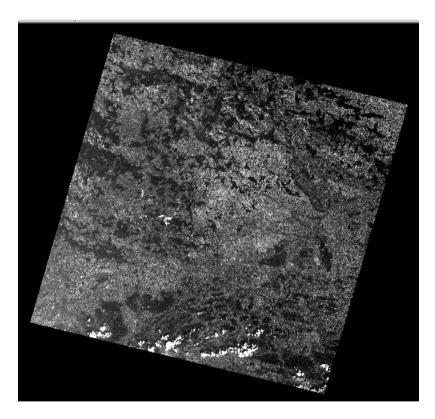
Image processing

- Pre-processing
- Image enhancement
- Data extraction

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Mean=45.49 Std.Dev=41.22	۸										
Median=76 Pred=0 (31149084)											
0.0% int= 0:255 0.5% int= 60:194											
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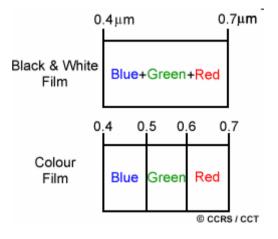
Histogram stretching

📓 Histogram "tm11" - TableHi	stogra	m(tm1_1.mpr) -	ILWIS								
<u>File E</u> dit <u>C</u> olumns <u>R</u> ecords <u>V</u> iew	<u>H</u> elp										
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🙀 Graph		1900000	1			1	67	1342967	1.84	3.22	36559:
Аьс							68	1548798	2.13	3.71	38108'
Legend		1800000	n i				69	1686122	2.31	4.04	39794;
tm1_1 - TableHistogram(tm1_		1700000				1	70	1741674	2.39	4.18	41536
X-Axis		1600000					71	1727613	2.37	4.14	43264
Abc		1500000					72	1694603	2.33	4.06	44958
		1400000-					73	1656567	2.27	3.97	46615:
117	<u>0</u>	1300000					74	1629282	2.24	3.91	48244.
Y-Axis Left	pixel	1200000-		.			75	1603456	2.20	3.84	49848
A⊳⊂ Number of pixels	l j	1100000		ll n	· · · · · · · · · · · · · · · · · · ·		76	1574039	2.16	3.77	51422
		1000000-		III n			77	1548641	2.13	3.71	52970'
2000000	Number	900000		h			78	1515138	2.08	3.63	54485
Y-Axis Right	L h	800000					79	1473512	2.02	3.53	55959:
Abc	Z	700000		IIIIIh			80	1419415	1.95	3.40	57378
		600000			L		81	1348689	1.85	3.23	58727
		500000-			L		82	1271567	1.75	3.05	599991
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innix - TableHistogram(tm1		100000-			111111111111	I II	Max	31149084	42.75	4.18	72860'
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Show record 80 in record view											11.

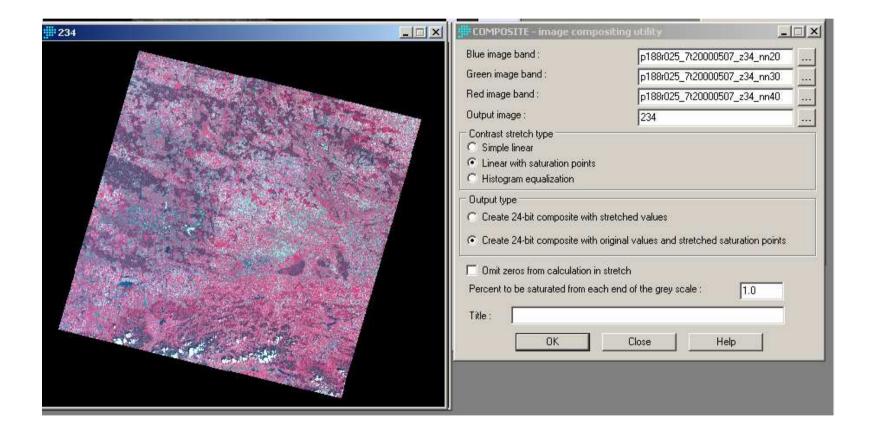
44

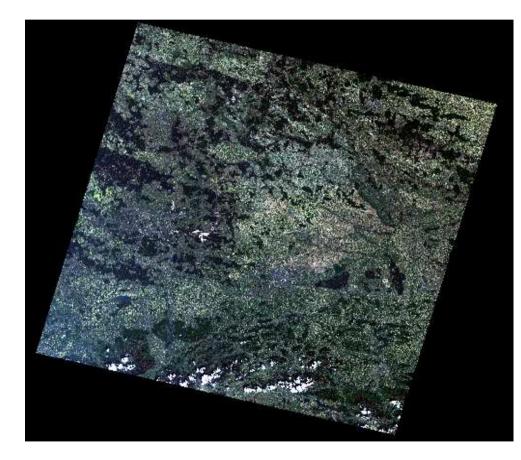
Histogram stretching

- Linear
- With saturation
- Histogram equalization

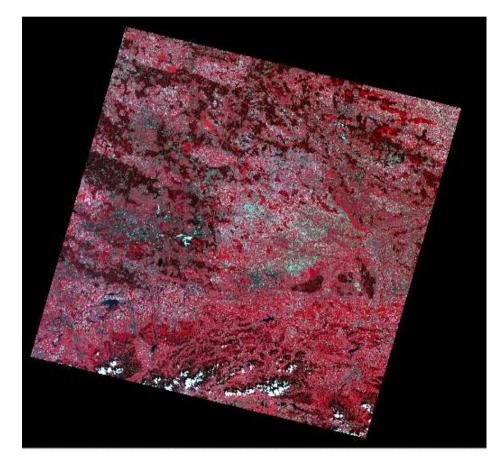


Blue image band : Green image band : Red image band : Output image :	p188r025_7t20000507_z34_nn10 p188r025_7t20000507_z34_nn20 p188r025_7t20000507_z34_nn30 123
Contrast stretch type C Simple linear Linear with saturation points Histogram equalization Output type C Create 24-bit composite with	stretched values original values and stretched saturation points





🔜 Color Composite		×
☑ <u>2</u> 4 bit ◎ <u>H</u> GB ○ <u>H</u> SI	 ☑ Linear Stretching ○ Histogram Equalization ☑ Percentage 	
<u>R</u> ed Band <u>G</u> reen Band <u>B</u> lue Band	Image: mage: mage	
Output Raster Map		



8	📑 Color Composite		×
	 ☑ <u>2</u>4 bit ④ <u>B</u>GB ○ <u>H</u>SI 	 Linear Stretching Histogram Equalization Percentage 	
	<u>R</u> ed Band <u>G</u> reen Band <u>B</u> lue Band	Image: mage 1.00 Image: mage 1.00 Image: mage 1.00 Image: mage 1.00	
	Output Raster Map	<u>S</u> how <u>D</u> efine Cancel Help	

